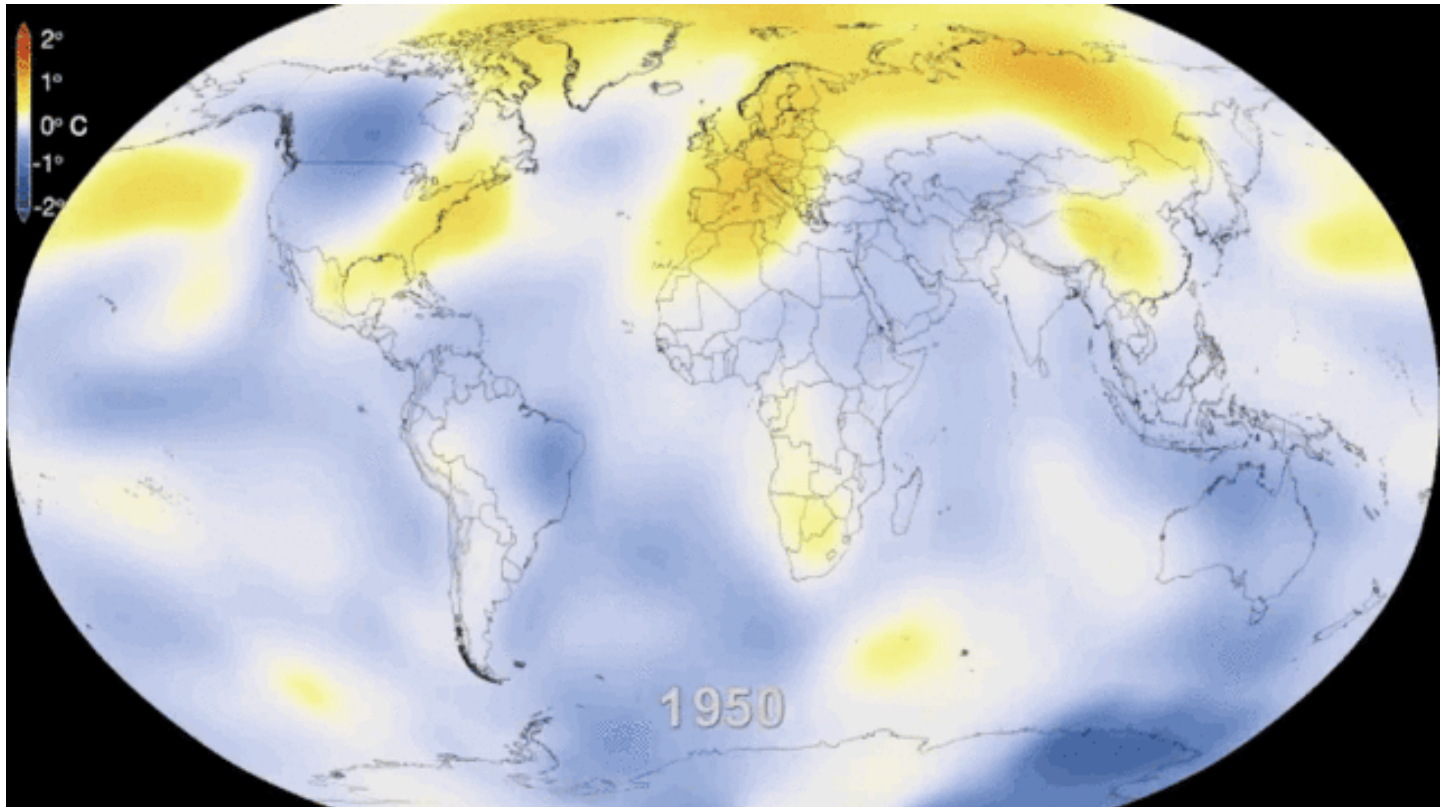


Variability of solar activity and climate impacts over the last centuries

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The primary energy source for the Earth's surface is the solar energy flux. While the flux received at the Earth's surface at a given location varies considerably, especially seasonally, the flux radiated by the Sun is relatively constant. However, variations in solar activity have been known for a very long time, and the hypothesis that this solar activity can affect our climate is an old one. What do we know about this solar activity over the last millenium; what activity can influence the climate; which influences are discernible?

1. The Sun, a variable star

Insolation, the flow of solar energy received on the Earth's surface, varies mainly because of the Earth's movements in its orbit, which causes daytime and seasonal alternation. On the contrary, **irradiance**, the flow of energy radiated by the Sun, has long been considered constant, hence its name "solar constant". It was not until the extremely accurate measurements obtained by satellite that **this "constant"** was shown to be **variable, but these variations are very small** (in the order of 0.1% on the time scales at stake here).

The main indicator of the Sun's own activity is the **sunspots**, these episodic dark spots on the Sun's surface. These spots had been known for a very long time, but it was the diffusion of the telescope at the very beginning of the 17th century that made it possible to count these spots regularly [\[1\]](#). These counts gradually revealed a very marked cycle with a 11-year period as well as variations on a century scale.

More recently, other expressions of this solar activity have been highlighted, including solar flares, extremely short phenomena

(a few hours at most) of radiation ejection and high-energy particles, which are responsible for the northern lights (read [Solar energies](#)).

1.1. The 11-year cycle (known as the Schwabe cycle)

The convective and magnetic activity of the outermost layers of the Sun, the photosphere, chromosphere and corona, shows marked and cyclical variations in its activity. When activity is strong, warmer and more radiant areas, the faculae, are formed on the visible 'surface' of the Sun (photosphere) accompanied by cooler and less radiant areas, the sunspots. These zones have an existence of the order of the week, and accompany the rotation of the Sun in 27 days. The solar atmosphere (chromosphere and corona) is also warmer and more radiant all over the Sun. Flashes are triggered more frequently from the corona.

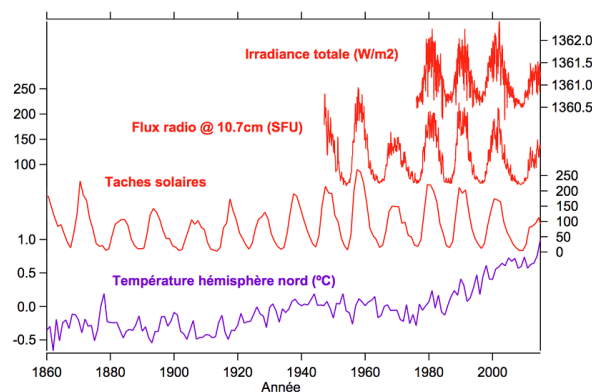


Figure 1. Changes in solar activity and surface temperature of the Earth over approximately the past 150 years. Radio flux measures the magnetic activity of the sun. [Sources: PMOD/WRC for irradiance; Hadley Centre/UEA CRU for temperature (HadCRUT4 version); Space Weather Centre of Canada for radio flux at 10.7 cm; Clette et al. 2014 (see ref.[1]) (WDC/SILSO) for annual sunspots]. [© G. Delaygue]

The sunspots, the most visible indication of this activity, have been counted since the beginning of the 17th century (around 1610) with the development of the astronomical telescope. An almost continuous series, annual to daily, could be reconstructed from counting archives [1]. This series of sunspot counts shows a very marked cyclicity at 11 years, with cycles varying from 9 to 13 years (Figures 1 and 2).

Satellite radiometers have been accurately measuring the variations in solar irradiance associated with this activity since about 1978. These measurements indicate a variation of 1 W/m², or **about 0.1%**, between the maximum and minimum of the total solar irradiance (Figure 1). These variations are stronger, relatively, in the ultraviolet wavelengths, in the order of 1 to 10%, even if their precise values remain debated [2].

1.2. Secular variations in solar activity ("Gleißberg" variations)

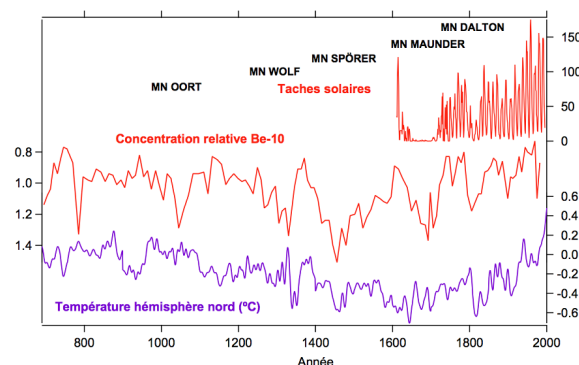


Figure 2. Solar activity (sunspots and cosmogenic isotope Be-10) and surface temperature of the Earth over the last 1200 years. The solar minima, in black, correspond to periods with few spots and high concentrations of Be-10. [Sources: Hoyt & Schatten (1997) for sunspot groups (distributed by NOAA/NCDC); Mann et al (2009) for northern hemisphere temperature (anomaly); Delaygue & Bard (2011) for relative variations of the Be-10 isotope in ice (inverted scale)]. [© G. Delaygue]

Sunspot count series show that the maximums of the 11-year cycles have fluctuated over the past 4 centuries, while the minimums have remained stable, with almost no sunspots. We thus observe periods characterized by very high maxima, periods that have been called "**Great maxima**". We are emerging from the "Great Modern Maximum", which culminated in the 1960s and 1970s. Periods of very low maxima, with very few spots, have been called "Great minima". Over the period covered by the stain count series, the most important is the **Maunder Great Minimum**, from about 1645 to 1715, during which very few spots were observed. These fluctuations in the secular time scale are often referred to as Gleißberg cycles (Figure 2).

While the number of sunspots is a good quantitative indicator of solar activity, it has the important bias of being levelled from below, with little or no sunspots during the Great Minima. However, the period covered by satellites does not include a Great Minimum: therefore, we do not know the minimum values of the irradiance, nor its total amplitude at this century time scale.

1.3. Secular to millennium variability: indirect indices of solar activity

Before these direct records of solar activity, **indirect indices** ('proxies') must be **used** to reconstruct it. The most reliable are the cosmogenic isotopes, Carbon-14 and Beryllium-10, whose production in the Earth's upper atmosphere is modulated by the solar magnetic field (Figure 2). These isotopes are recorded in tree rings and the ice on polar ice caps and mountain glaciers. Their concentration is interpreted in terms of solar activity, but only relative variations can be reconstructed. To obtain absolute variations, for example of total solar irradiance (in W/m^2), it is necessary to **calibrate these concentrations** with independently known variations, either measurements (but restricted to a very short and recent period) or estimations based on physical models of solar activity [3].

However, these indirect indices remain relatively imperfect, both in terms of reliability and temporal resolution. In particular, the 11-year solar cycle remains very poorly resolved in the long records of these proxies, even if recent studies have highlighted it over the last millennium [4].

Over the past millennium, these proxies show that solar activity was low not only during the Maunder Great Minimum but also, and perhaps even lower, during the Spörer Great Minimum in the 15th century. The modern Grand Maximum, at the end of the 20th century, seems a little higher than the medieval Grand Maximum of the 8th century (Figure 2).

2. Impacts of the 11-year solar cycle on the Earth's climate

Following the identification of 11-year solar cycles, countless studies have sought to find correlations with climate variations or other parameters [5]. Few of these correlations were significant and withstood serious analysis [6].

2.1. Impacts on the Earth's surface

It now seems clear that the variations in irradiance due to the 11-year cycles are too small and too fast to generate significant temperature variations at the Earth's surface, given its thermal inertia. In fact, the maximum amplitude of the irradiance, of the order of $1 W/m^2$ during a solar cycle, modifies the solar flux absorbed by the surface by about $0.14 W/m^2$, taking into account the geometric dilution factor on the terrestrial sphere (factor 1/4), the global albedo (0.3), and the absorption by the atmosphere (about 20%). The impact on a hemispherical or global average of the surface temperature seems just detectable, in the order of $0.1^\circ C$ per cycle (link : Forcing and retroactive effects; climate sensitivity).

However, a persistent correlation has been shown between these solar cycles and surface pressure and temperature in some regions of the Pacific [7]. It also appears that this solar cycle interacts with oscillations within the climate system, including those of the same decadal time scale, the Pacific decadal Oscillation and the North Atlantic Oscillation [8]. At the regional scale, solar activity could favour some synoptic regimes, particularly in winter, blocking regimes that favour the advection of polar cold air over Western Europe [9].

2.2. Impacts on the stratosphere

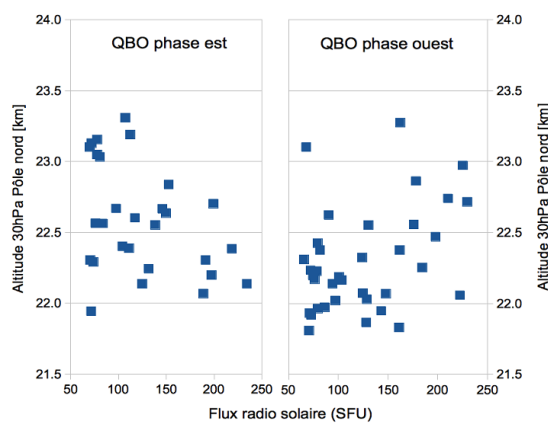


Figure 3. The impact of solar activity on atmospheric pressure over the North Pole depends on the QBO phase. Relationship between the altitude of the 30hPa geopotential at the North Pole (mean of January-February; NCEP reanalysis) and the solar radio flux at 10.7cm, over the period 1948-2015, with a distinction of the direction of the low stratospheric jet winds (QBO phases). [© G. Delaygue]

Unlike the surface, this **11-year-old solar signal is well marked in the stratosphere**. Variations in temperature, geopotential and meridian circulation (known as Brewer-Dobson's circulation) are well documented (Figure 3). The mechanism involved, tested using models describing physical and chemical processes [10], is the strong absorption of UV radiation by the ozone cycle, which warms the air and modifies the meridian temperature gradients. The relative variations in UV radiation during a solar cycle are large compared to the variations in total irradiance (10 to 100 times larger, relatively). These impacts on the stratosphere affect the intensity of the Arctic polar vortex, especially in winter, which in turn affects the lower troposphere at high and mid-latitudes. These impacts of the solar cycle are modulated by the direction of circulation of the stratospheric jet stream, a direction that is reversed about every two years ('quasi-biennial oscillation', QBO) [11] (Figure 3).

2.3. Possible impacts on cloud formation

Solar activity modulates not only the radiation of the Sun but also the flow of energy particles that penetrate the Earth's atmosphere. This flow is primarily formed by the galactic cosmic radiation, deflected by the solar magnetic field (which varies with the activity of the Sun). Energetic particles are also emitted by the Sun itself, particularly during eruptions. These energetic particles interact with molecules in the atmosphere to form ions. The hypothesis that this ionization has a climatic impact dates back to more than 50 years [12], but remains highly debated, both theoretically and in observations. The mechanism most often advanced is that these ions would promote the formation of condensation nuclei necessary for cloud formation.

The role of clouds on heat exchange in the climate system is fundamental, via radiative fluxes, in visible and infrared wavelengths, but also latent heat fluxes. Solar modulation of the flow of energetic particles has a very limited direct climatic impact, but a very strong amplification could exist via this control on clouds. However, the various proposed correlations between energetic particle fluxes and cloud cover remain highly debated [13].

3. Impacts of secular variability on climate: the last millennium

The Maunder Great Solar Minimum (~1645-1715), well known by sunspots, took place during a period described by historians in Europe as the "Little Ice Age" [14]. However, this cool period remains poorly defined in time, probably because the decadal variability was strong compared to the climate change, and also because its definition depends on the considered criteria and regions. In its wide acceptance, the Little Ice Age extends from the 15th century to the middle of the 19th century, an end marked by the massive retreat of alpine glaciers.

Interpretation of historical and natural records (tree rings, in particular) in terms of local climatic conditions (temperature and/or precipitation) confirmed that the Little Ice Age was a detectable phenomenon on the scale of the Northern Hemisphere, but with a small amplitude (about 0.2°C) compared to regional and temporal variability (i.e., colder and warmer decades).

Climate models are used to test the respective importance of different factors of climate change, including solar variability. These factors are taken into account by the radiative imbalance they are responsible for ('climate forcing', [link: point 1.3]). Recent studies agree with a small but detectable contribution of the solar activity to surface temperature variations from regional to global scales. However, the contribution of strong volcanic eruptions to these temperature variations has been revised upwards by these studies, including at decadal and even centennial scales [15].

4. Conclusion

It therefore seems likely that solar activity has had more pronounced climate impacts at the regional scale than at the global

scale, at the time scales considered here. Better reconstructions and simulations of these regional climate variations, and independently of climate forcings, are needed to understand their relationships. These studies of the last few centuries make it possible to consider the climatic impact of the Sun over the next few centuries, particularly with the hypothesis of a return to a very low level: this impact should remain low on a global scale compared to those of anthropogenic forcings (greenhouse gases, albedo), but at a regional scale it will continue to interact with climate variability [16].

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Cover picture. Increase in global temperatures from 1950 to the end of 2013. [Source: NASA (Public domain), via Wikimedia Commons]

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